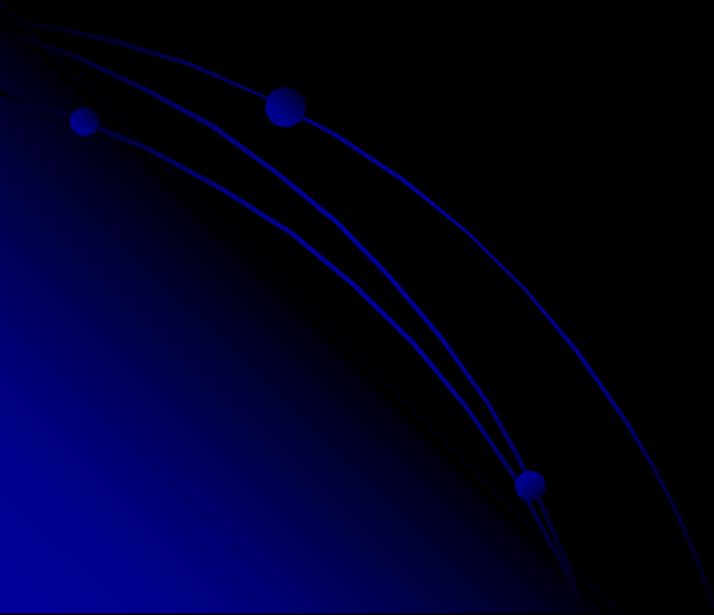


# Sedimentary Rocks and Processes



# Sedimentary Processes

- **Weathering**
  - Breakdown of pre-existing rock by physical and chemical processes
- **Transport**
  - Movement of sediments from environments of relatively high potential energy
- **Deposition**
  - Settling of sediments in environments of relatively low potential energy
- **Diagenesis**
  - Chemical changes that occur after sediments are buried

# Sedimentary Processes

- **Weathering**

- **Physical Weathering**

- Mechanical breaking of rock into smaller pieces (clasts)
- No new minerals formed

- **Chemical Weathering**

- Reaction of rocks with air or water
- New minerals formed at the expense of less stable minerals

- **Transport**

- **Deposition**

- **Diagenesis**

# Physical Weathering

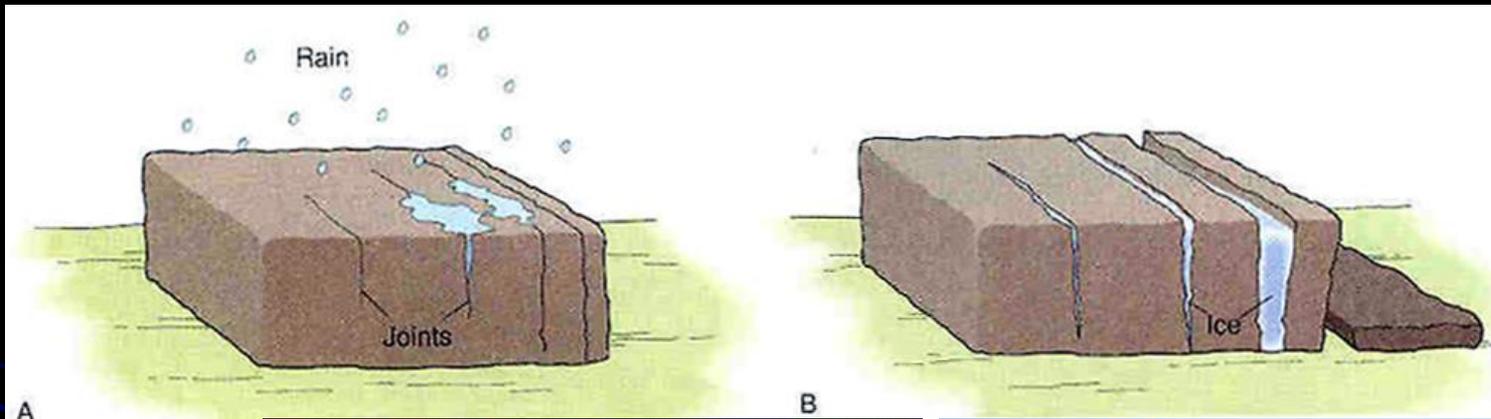
- **Abrasion**
  - Grinding/scratching by rock fragments (e.g., sandblasting)



# Physical Weathering

- **Frost Wedging**

- Ice has approximately 8% greater volume than liquid water
- Ice in cracks expands and creates an extensional stress



A

B

Frost Wedging



# Physical Weathering

- **Root Wedging**

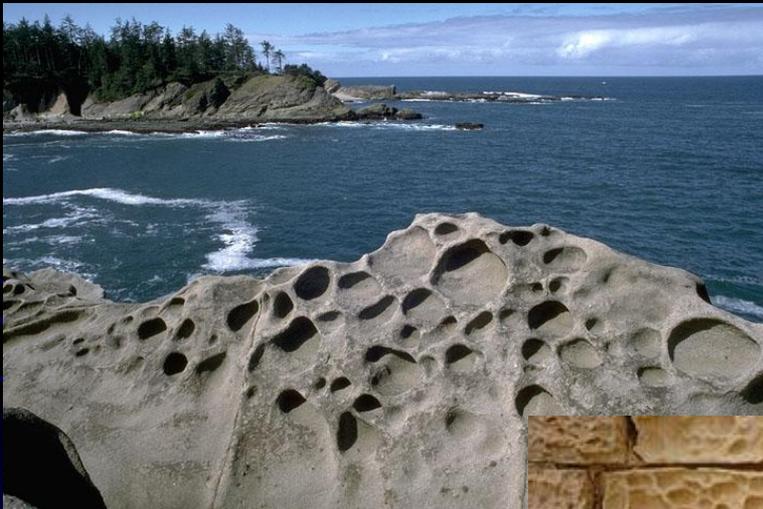
- Trees take root in existing cracks in rocks
- As tree grows, the expanding root pries rock apart



# Physical Weathering

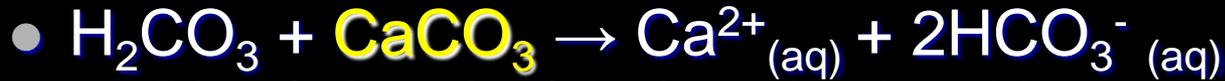
- **Salt Wedging**

- Salt crystals that precipitate in pores due to evaporation of saline solutions will expand when heated and create stress



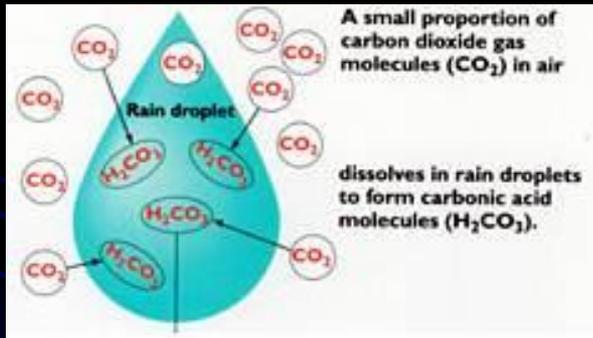
# Chemical Weathering

- Dissolution



- Dissolution of calcite in carbonic acid

- All rainwater is a weak carbonic acid solution



# Chemical Weathering

## ● Hydrolysis

- In hydrolysis a molecule is split into two parts by reacting with water. One product includes an  $\text{OH}^-$  from the water molecule, whereas  $\text{H}^+$  is added to another



- Hydrolysis of **olivine produces a solution** (mineral dissolves)



- Hydrolysis of **K-feldspar produces kaolinite** (clay product)

Silicic Acid  
(Dissolved Silica)



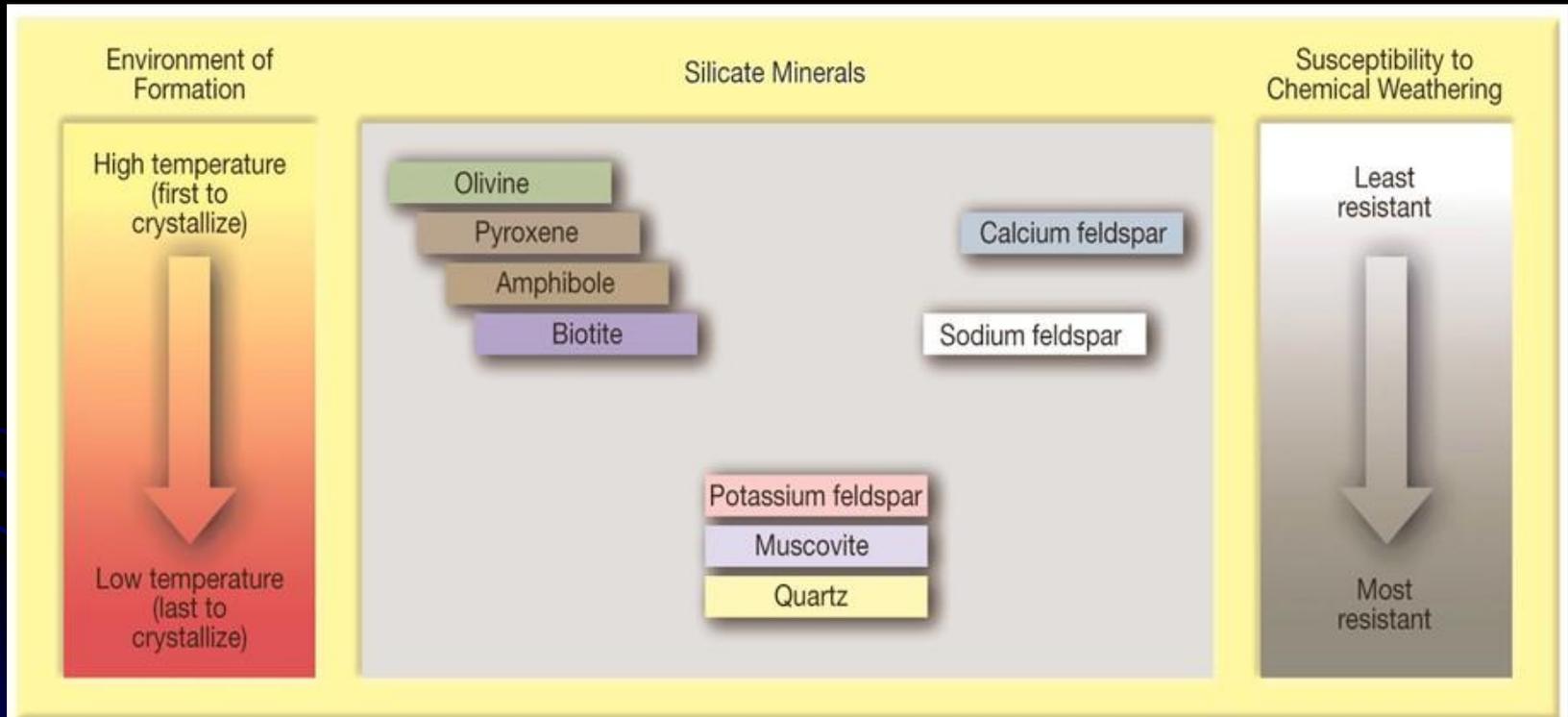
# Chemical Weathering

## ● Oxidation

- $4\text{FeS}_2 + 15\text{O}_2 + 22\text{H}_2\text{O} \rightarrow 4\text{FeO}(\text{OH}) + 8\text{SO}_4^{2-} + 16\text{H}^+$ 
  - Oxidation of pyrite to form limonite
- $4\text{FeSiO}_3(\text{s}) + \text{O}_2(\text{g}) + 8\text{H}_2\text{O} \rightarrow 2\text{Fe}_2\text{O}_3(\text{s}) + 4\text{H}_4\text{SiO}_4(\text{aq})$ 
  - Oxidation and hydrolysis of pyroxene to form hematite



# Bowen's Reaction Series and Chemical Weathering



# Chemical Weathering

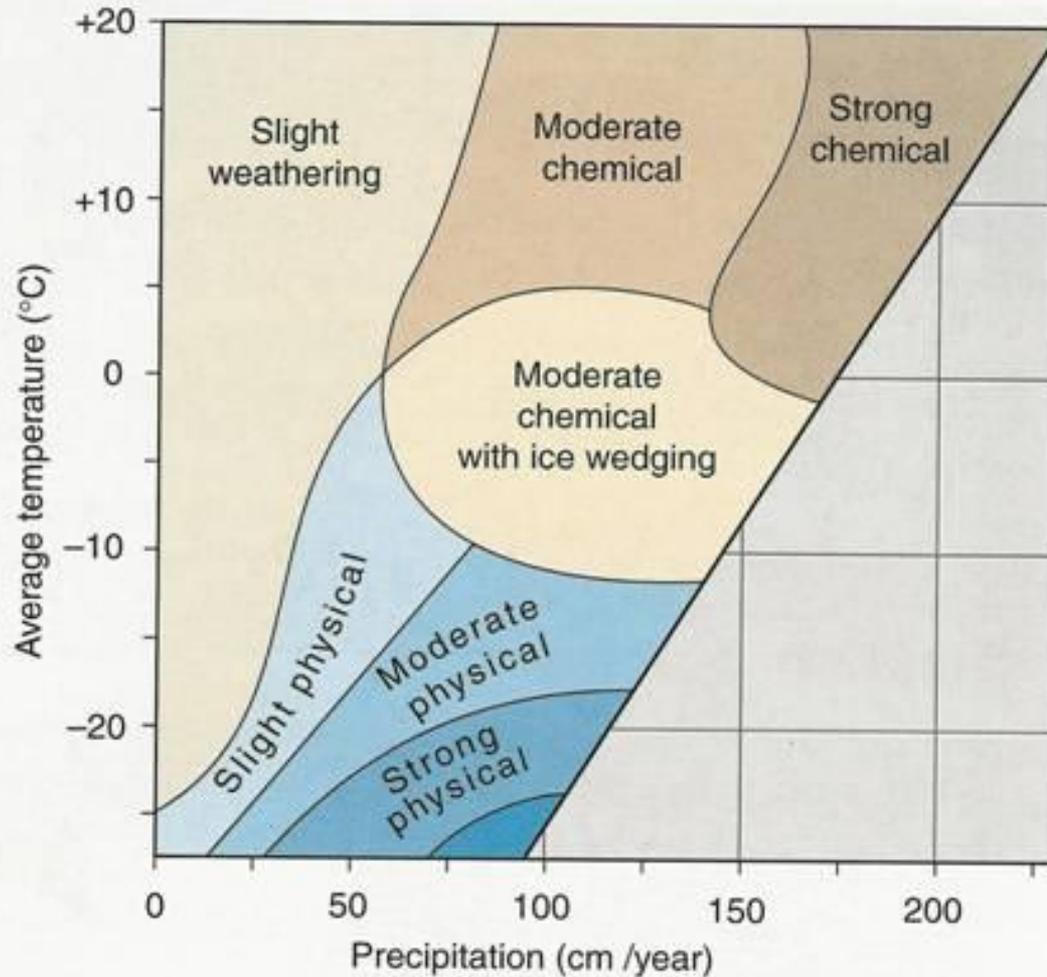


Start with feldspars and Fe-bearing Silicate minerals...



End with clays, limonite, hematite, and a loss of Si, Ca, Na, and K into solution

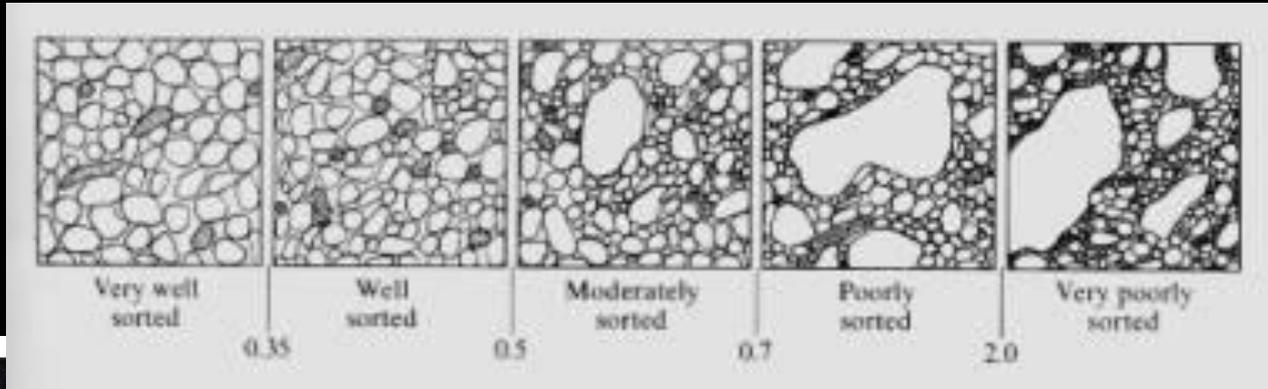
# Weathering and Climate



# Sedimentary Processes

- Weathering
- Transport
  - Movement of sediments from environments of relatively high potential energy
  - Moved by wind or water
    - Sorted and stratified
    - Relatively fine-grained (sand by air; pebbles by water)
  - Moved by glacier or gravity flow
    - Unsorted and unstratified
    - Clay to boulders
- Deposition
- Diagenesis

# Sorted vs Unsorted Sediments



Well-sorted sand



Poorly sorted sand

# Sorted vs Unsorted Sediments

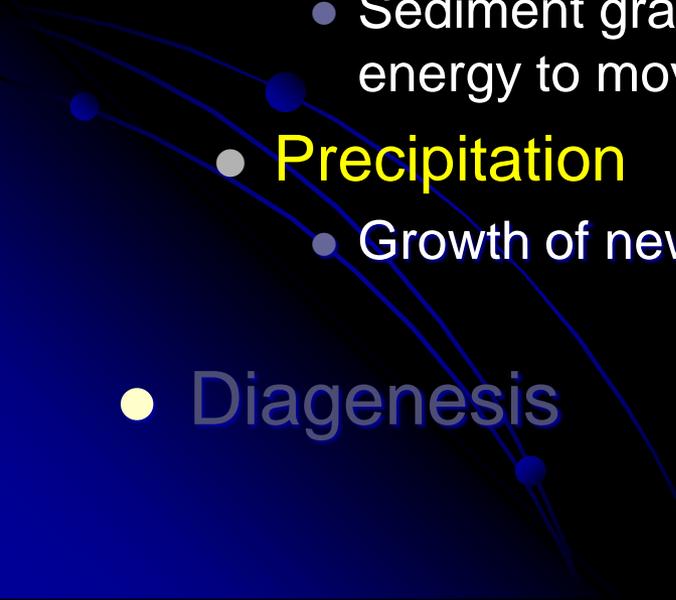


Well-Sorted Stream Deposits

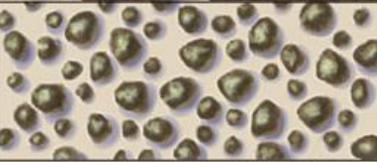
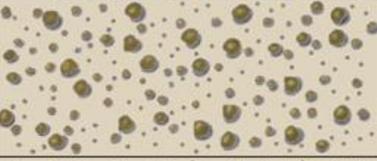


Poorly Sorted Glacial Deposits

# Sedimentary Processes

- Weathering
  - Transport
  - Deposition
    - Settling
      - Sediment grains stop moving when there is insufficient energy to move them farther
    - Precipitation
      - Growth of new minerals from oversaturated water
  - Diagenesis
- 

# Sediment Classification by Grain Size

A. Grain size		
"Gravel" > 2mm	Pebbles 4–64 mm	
	Granules 2–4 mm	
	Coarse sand 0.5–2 mm	
	Medium sand 0.25–0.5 mm	
	Fine sand 0.06–0.25 mm	
	Silt 0.004–0.06 mm	
	Clay < 0.004 mm	

Conglomerate

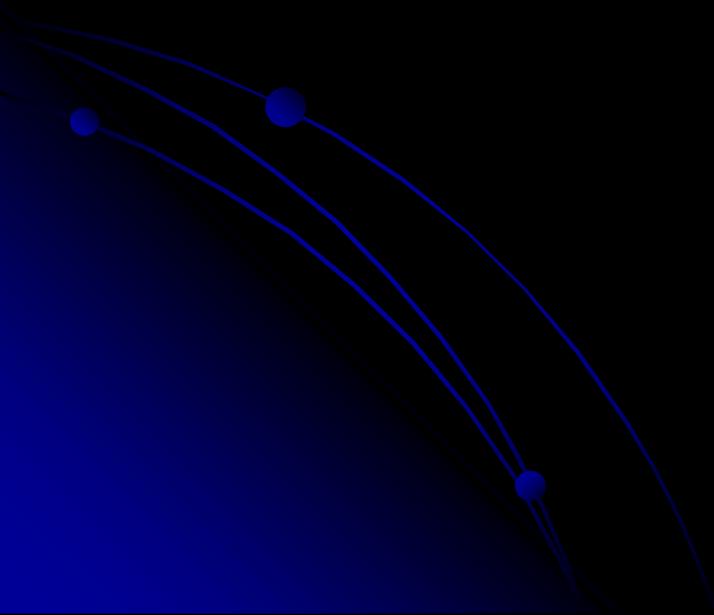


Sandstone

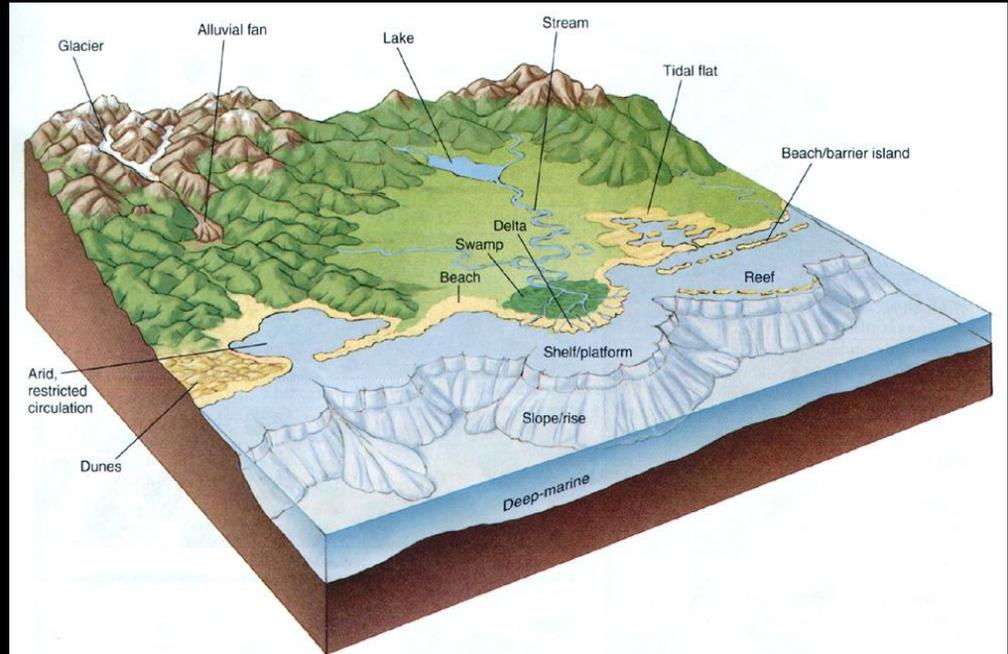
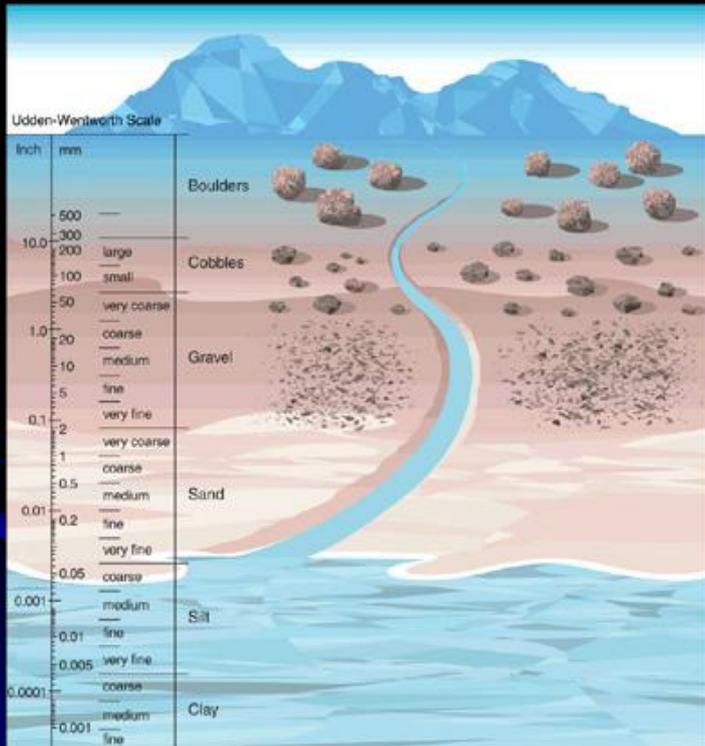


Siltstone

Mudstone/Shale



# Grain Size and Depositional Environments



**FIGURE 4.2**  
 Typical sedimentary depositional environments.  
 (Adapted from Jones, 2001: Laboratory Manual for Physical Geology, 3rd edition.)

# Deposition by Precipitation (Chemical Sediments)

- Normal seawater has a salinity of 35g/kg
- Only 6 ions comprise 99% of the ions dissolved in seawater:
  - Cl<sup>-</sup> 55.04 wt%
  - Na<sup>+</sup> 30.61 wt%
  - SO<sub>4</sub><sup>2-</sup> 7.68 wt%
  - Mg<sup>2+</sup> 3.69 wt%
  - Ca<sup>2+</sup> 1.16 wt%
  - K<sup>+</sup> 1.10 wt%
- Marine precipitates will be composed of these ions.



# Deposition by Precipitation (Chemical Sediments)

Mineral	Composition	Solubility (mol/l)
Calcite	$\text{CaCO}_3$	0.00014
Gypsum	$\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$	0.0154
Magnesite	$\text{MgCO}_3$	0.001
Halite	$\text{NaCl}$	6.15
-----	$\text{CaCl}_2 \cdot 6\text{H}_2\text{O}$	7.38

*Why is  $\text{CaCl}_2 \cdot 6\text{H}_2\text{O}$  not a naturally occurring evaporite mineral?*

# Chemical Sediments

- Limestone are the most common chemical sediment
  - Common in shallow, tropical, marine environments
- Evaporites (deposits of gypsum and halite) are restricted to hot, arid environments where evaporation rates are high

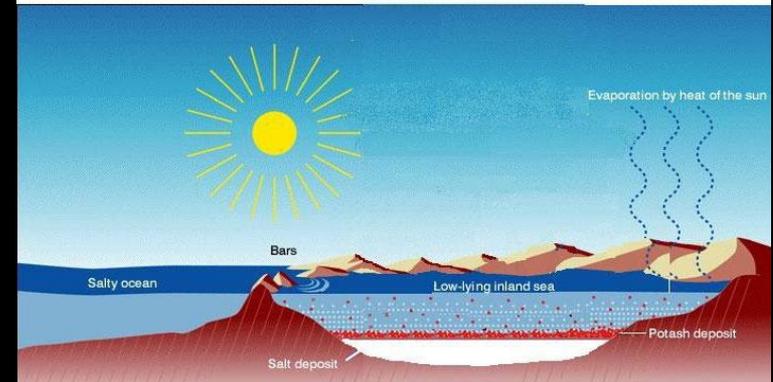
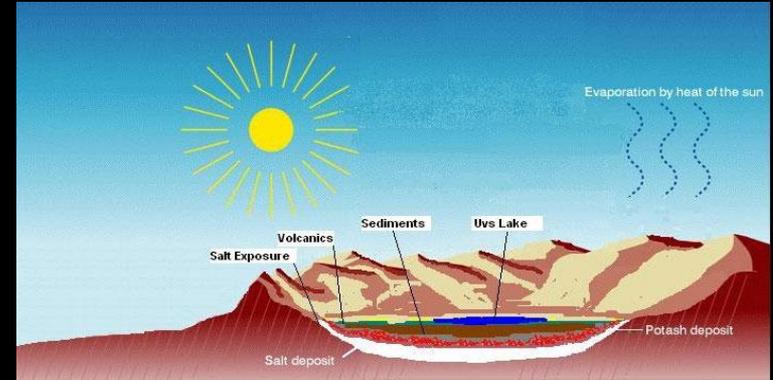
# Limestone Reef and Lagoon



# Fossiliferous Limestone



# Evaporite Deposits



# Diagenesis: Cementation

- Precipitation of minerals from solutions in pore water
- Composition of ground water varies greatly depending on such variables as bedrock composition and rainfall
  - *Calcite and dolomite might develop in hard-water areas*
  - *Clays and quartz might develop in soft-water areas*
  - *Limonite and hematite may occur in environments rich in iron (terrestrial)*
  - *Pyrite may occur in rocks where the pore waters were anoxic*

